Turbulence structure in a fractal forest under varying atmospheric conditions

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Outline

I. Motivation

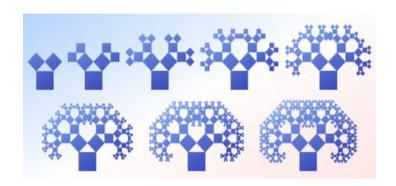
II. Method

- a) Ensemble of Fractal Trees
- b) Heated Immersed Boundaries
- c) Resolved Flow: 100m to 5cm

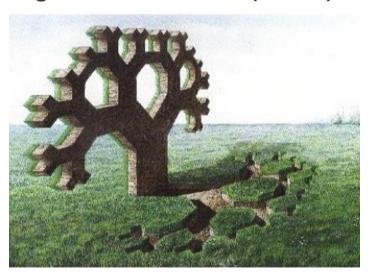
III.Results

- a) Plant Scale Approach
- b) Coherent Structures

IV. Conclusions



Pythagoras tree (below) and algorithm described (above)



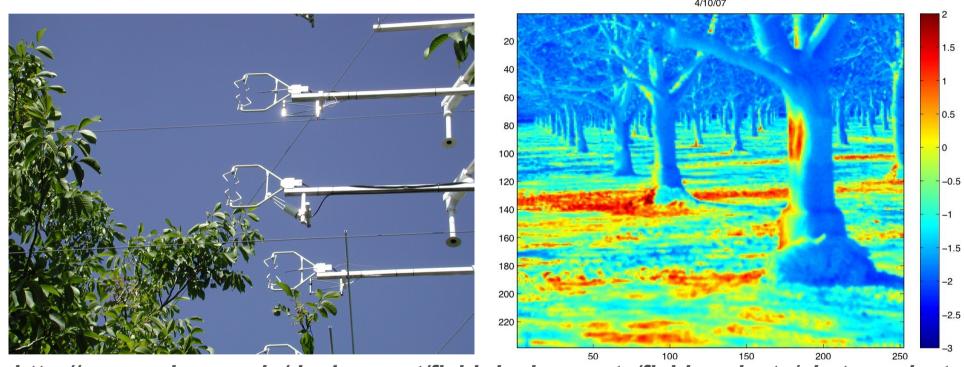






Motivation

In **forests** measurements are taken at the microscale from meters **down to a few cm**.



http://www.eol.ucar.edu/deployment/field-deployments/field-projects/chats-project

Eddy Covariance Sensors in a Walnut Canopy (left) and infrared image of heterogeneously heated trees and trunk space (right)

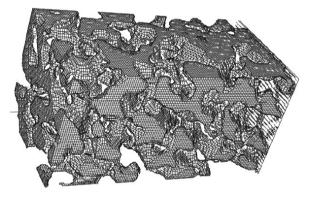






Field Scale Approach

"Forest as a **porous body** of horizontally uniform **leaf** area density: LAD(z) with constant drag coefficient c_D " (Shaw & Schumann 1992)



 $\boldsymbol{F}_{\boldsymbol{D}} = -c_D LAD(z)|\boldsymbol{u}|\boldsymbol{u}$

Field-scale simulations, where resolution is of O(1 m)

- Shaw & Schumann (1992)
- Dupont & Brunet (2009)
- Finnigan et al. (2009)

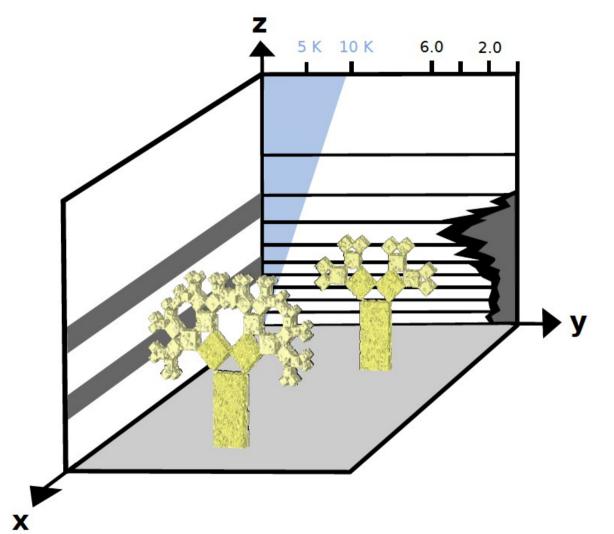
Is it possible to resolve the turbulence structure correct over this wide range of scales by state-of-the art multiscale numerical simulations?







Plant Scale Approach



- Ensemble of 16 trees, vary in a Gaussian way: height, fractality, position, scale-dependent porosity
- Thermal Stability of ambient air (Shaw et al. 1988, Gao et al. 1989)
- Heated Tree Crown (EAGLE, CHATS: 3K)
- Vertically Stretched grid across surface layer (100m,10m,10cm)





EULAG, LES with Immersed Boundaries

$$\nabla \cdot \mathbf{v} = 0$$

$$\frac{d\mathbf{v}}{dt} = \nabla \frac{\rho'}{\rho_b} - \mathbf{g} \frac{\theta'}{\theta_b} + \mathbf{D}^{\mathbf{v}} - \beta(\mathbf{v} - \mathbf{v_F})$$

$$\frac{d\theta'}{dt} = \mathbf{v} \cdot \nabla \theta_e + \mathbf{D}^{\theta} - \beta(\theta - \theta_F)$$

$$\frac{d\mathbf{e}}{dt} = \mathbf{S}(\mathbf{e}) - \beta(\mathbf{e} - \mathbf{e_F})$$

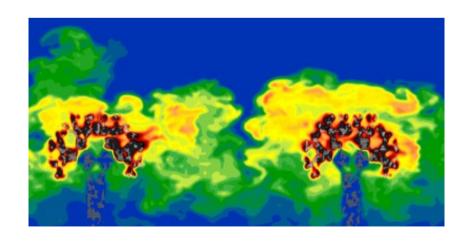
Boussinesq Approximation

 $\rho_{\rm b} = 1.025 \; {\rm kg/m^3}$

 $\Theta_{p} = 300 \text{ K}$

 $p_{h} = 1000 \text{ hPa}$

 $\Theta_{\rm F} = \Theta_{\rm e} + 3.15 \, \rm K$



Imrsb. w/ a prescribed temp. are an extension to the ones used for 'Building resolv. LES & comparison with windtunnel studies' (Smolarkiewicz et al. JCP 2007)

'EULAG, a computational model for multi-scale flows' (Prusa et al. 2008)







Experimental Setup

Stretched vertical coordinate

 $\Delta x = \Delta y = 5 \text{ cm}$

 $\Delta z = 12cm, ..., 12m$

Domainsize

Gridpoints 384 x 384 x 384

19.2 m x 19.2 m x 108 m

Periodic lateral boundaries

Timesteps

 $\Delta t = 0.002 s$

nt=180 000, Time=360s

Moving average:

Online statistics over last 5 min

Runs	N [1/s]	ΔT [K]	LAI	<i>U</i> [m/s]
1) neutral	0	0	2.8	2.8
2) n+heat	0	3.15	2.8	2.8
3) n+heat	0	3.15	1.9	2.8
4) stable	0.05	0	2.8	2.8
5) s+heat	0.05	3.15	2.8	2.8

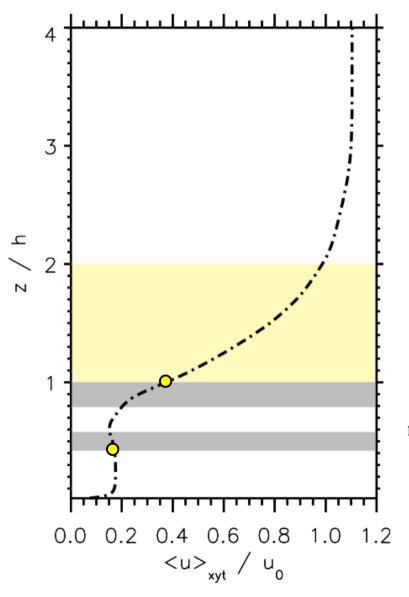
'Turbulence structure in a diabatically heated forest canopy composed of fractal Pythagoras trees' (Schröttle & Dörnbrack, TCFD 2012)







Velocity Profile



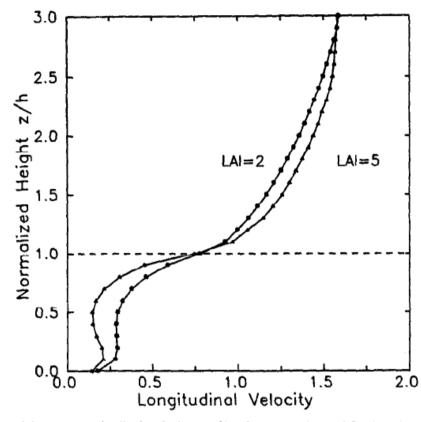


Fig. 3. Spatial mean longitudinal velocity profiles for two values of LAI under weakly unstable conditions. Velocity is normalized by the vertically averaged longitudinal velocity.



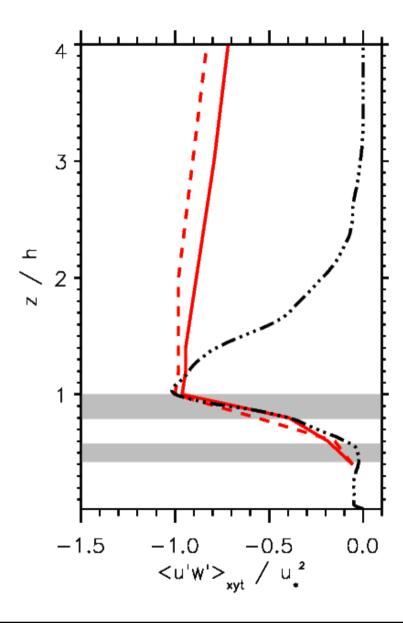
Shaw & Schumann, BLM 1992 Schröttle & Dörnbrack, 2012 Inflection Point Vorticity Thickness of IBL







Momentum Transport



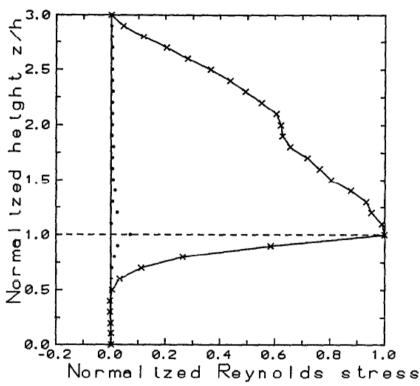


Fig. 4. Vertical profile of the spatial mean Reynolds stress for a LAI of 5 and weakly unstable conditions, and normalized to its value at the top of the canopy. The solid line is the sum of resolved and subgrid-scale components of the Reynolds stress. The dots are the SGS component.

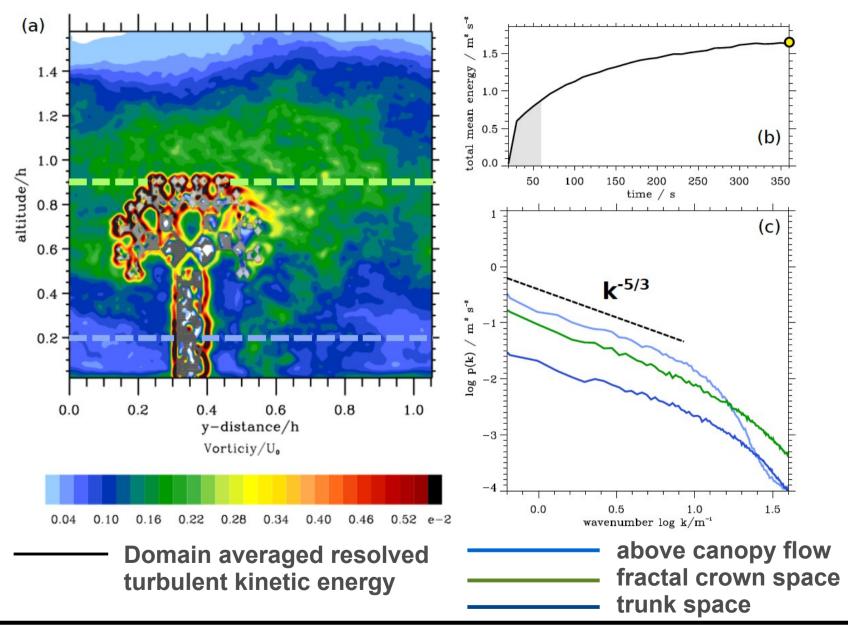
Shaw & Schumann, BLM 1992 Brunet Windtunnel, BLM 1994 Finnigan et al. LES, JFM 2009 Schröttle & Dörnbrack, 2012







Vorticity, Turbulent Kinetic Energy

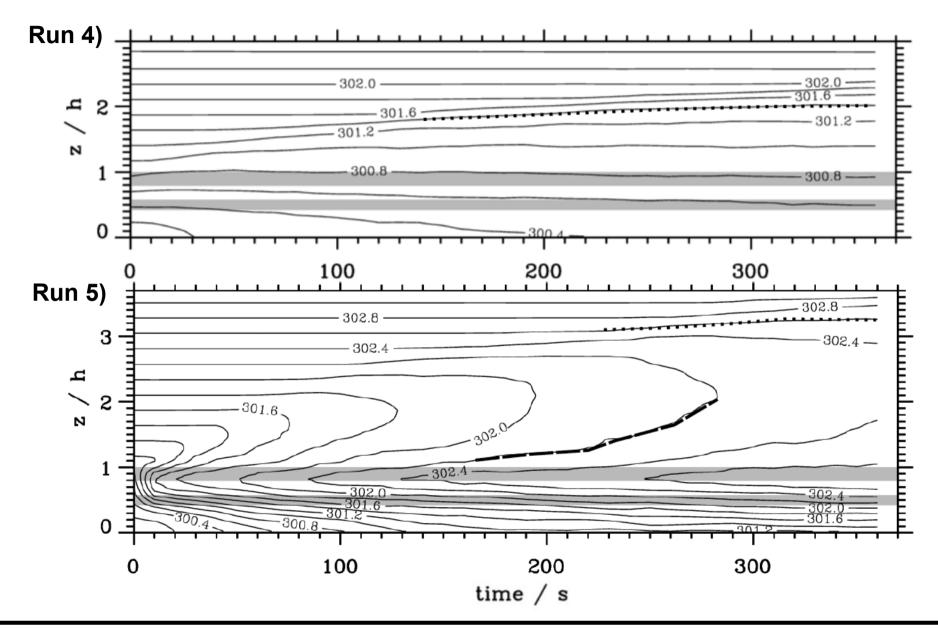








Stratification, Hovmoeller Diagrams

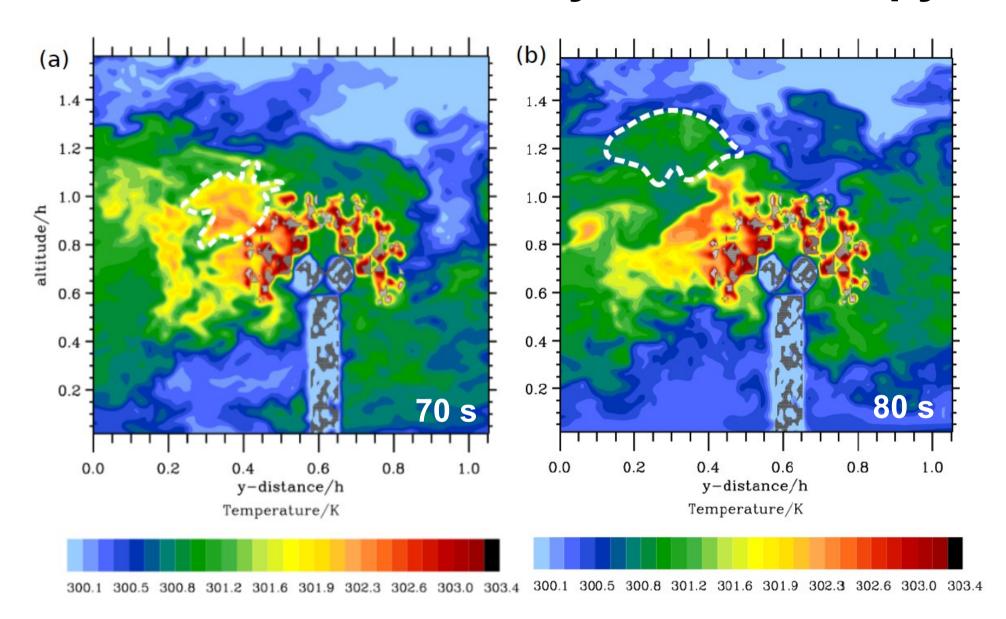








Thermals in diabatically heated canopy

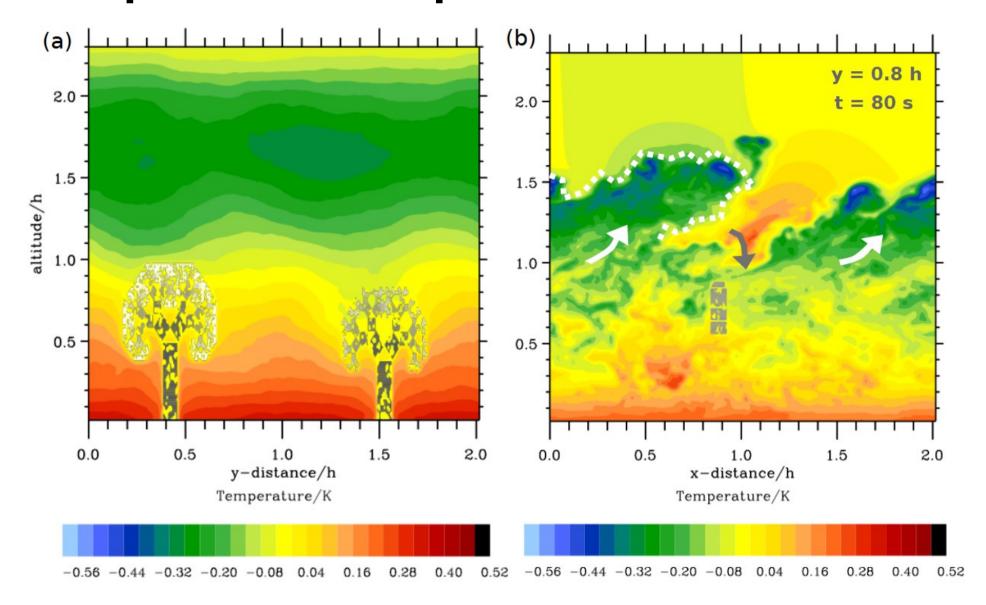








Temperature Ramps in stable conditions

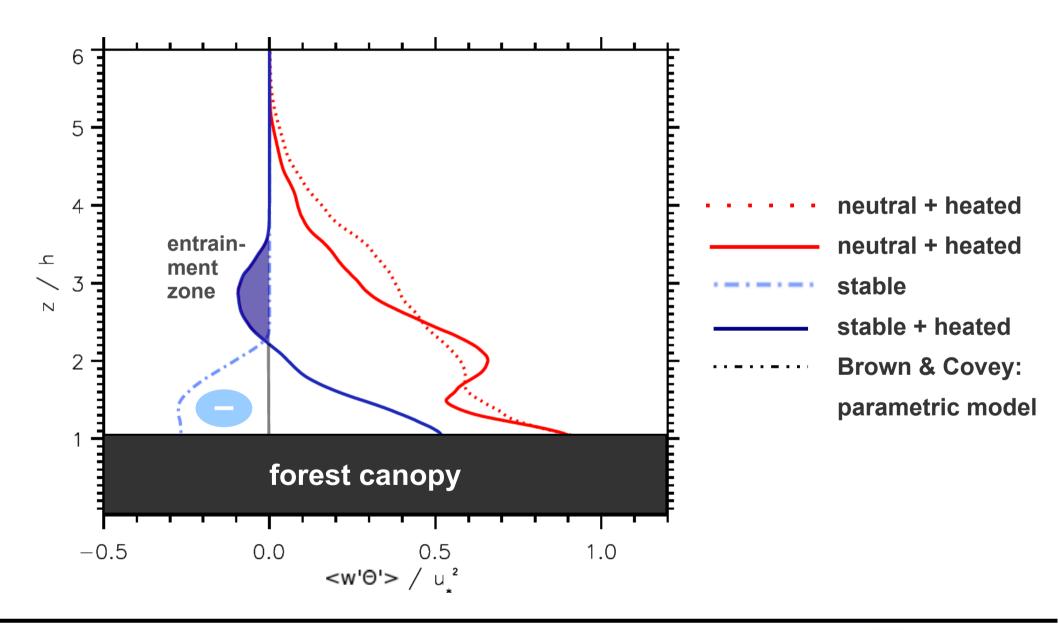








Heat Flux for varying Stratification

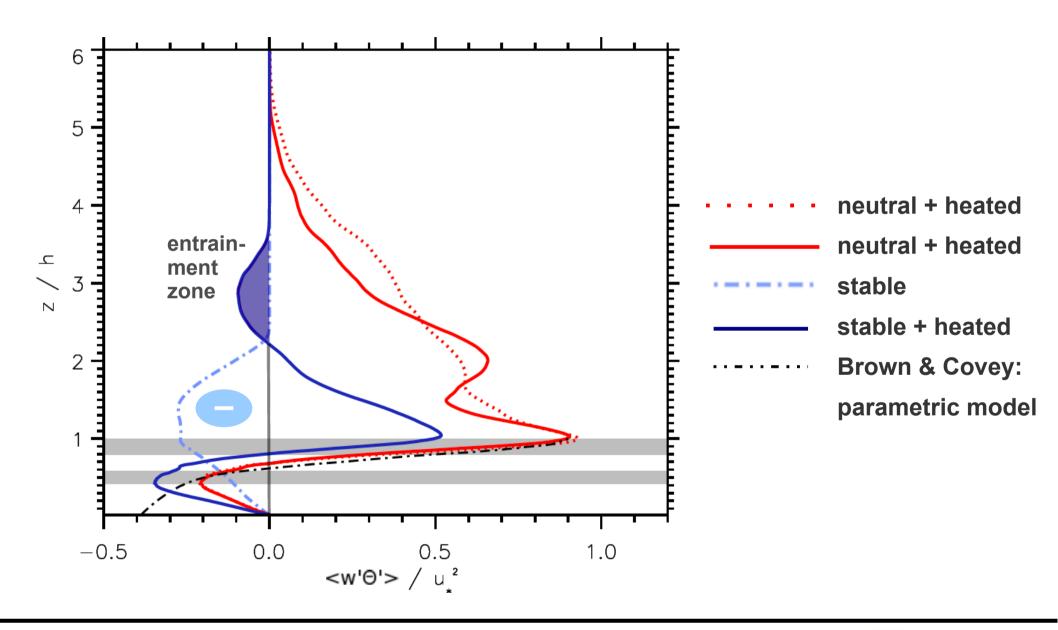








Heat Flux for varying Stratification









Conclusions

Outlook

METHOD

Applicability of **fractal approach** in LES: resolve flow, get **physically correct results** in **neutral** reference run (*Dupont & Brunet 2009, Finnigan et al. 2009, Shaw & Schumann 1992*).

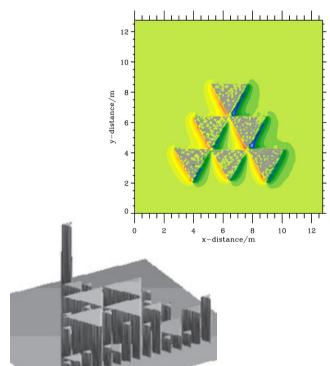
FLOW STRUCTURE & DIABATIC RESULTS

Characteristic power spectra for trunk-, crown- and above canopy flow occur. As in field experiments, we capture wake vortices inside the canopy (*Cava & Katul 2007*).

Coherent structures in diabatic runs are observed as in nature (*Gao et al. 1989*). We can simulate them in detail in our LES.

Thank you for your attention!





Master student **Sonja Gisinger**: *Sierpinski* city





